# Sediment Transport Modeling in GSSHA



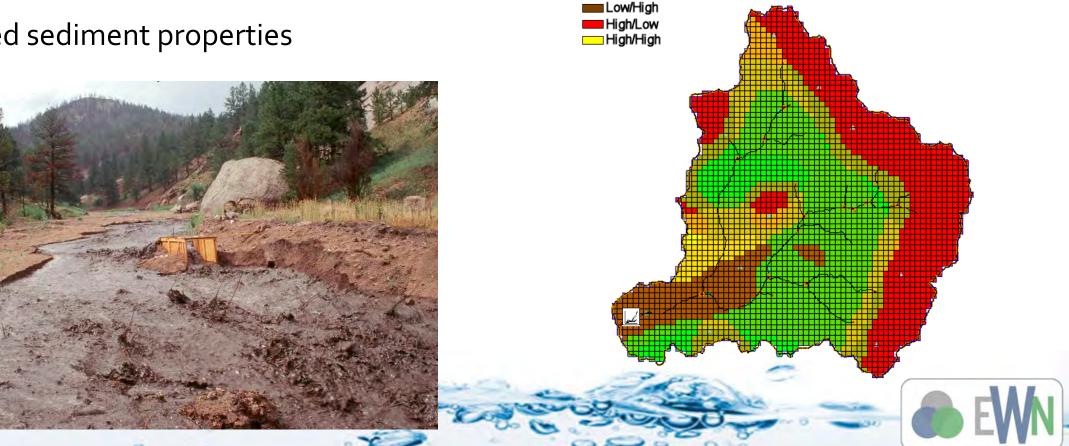


### **Sediment Transport Overview**

Erosion/Deposition

Low/Low

- Event based erosion and deposition model (not USLE-based)
  - Overland
  - Streams
- User-defined sediment properties



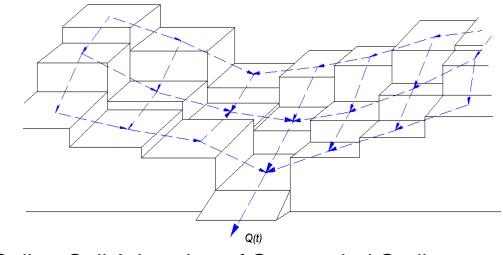
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- Simulation of sediment erosion from the overland plane
- Transport across overland plane including deposition and reentrainment
- Transfer into channels
- Channel routing of bed and wash loads
- Routing of sediments through reservoirs
- Landscape evolution with the capability to do multi-decadal simulations

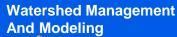


- Any number of sediment particles
  - Size
  - Specific gravity
- Sediment detachment
  - Raindrop impact
  - Overland flow limits
- Transport Capacity
  - Kilinc-Richardson
  - Engelund-Hansen
  - Multiple shear stress formulas
- Deposition



Cell to Cell Advection of Suspended Sediments







- Particles are detached by rainfall impact and overland shear.
- If sufficient transport capacity is available, the eroded particles are transported to adjacent cells by advection.
- Deposition is computed with trap efficiency.
- Transport and deposition by particle size and density.
- Sediments are accounted for using three storages:
  - Parent material
  - Deposited materials
  - Materials in suspension
- Each pool has its own distribution of particles.
- Deposited materials are assumed to erode first, then the parent material.
- Elevations can be adjusted for erosion/deposition.





#### **Factors Affecting Detachment and Transport Capacity**

- Soil particle size finer particles are more erodible.
- Cover conditions any type of vegetative or manmade cover protects the particles from detachment and transport.
- Management actions actions taken to prevent erosion, no-till, contour plowing, etc. lesson erosion.
- Erosive forces
  - Rainfall intensity
  - Discharge
  - Land surface or friction slope
  - Shear stress





- Detachment of particles by rainfall is a function of momentum, which is a function of rainfall intensity.
- Rainfall detachment (kg m<sup>-2</sup> s<sup>-1</sup>) is calculated as the product of several factors related to:
  - K<sub>I</sub> soil erodiability
  - C<sub>w</sub> correction for water depth
  - C<sub>G</sub> vegetative canopy cover factor
  - C<sub>i</sub> cover-management factor, and
  - M<sub>R</sub> rainfall momentum squared.

## $D_R = K_I C_w C_G C_i M_R$





- Overland flow detaches particles by exerting a shear stress that breaks particle bonds.
- Erosion occurs in rills.
- Within a cell overland flow and rill erosion are assumed uniform, so that calculations are for gross rill erosion.
- Detachment capacity rate (kg m-2·s-1)

$$D_{c} = a(\tau - \tau_{c})_{r}^{b}(1 - G/T_{c})$$

- *a* and *b* are empirical coefficients,
- $\tau$  = the flow shear stress (Pa),
- τ<sub>cr</sub> = the critical shear stress,
- $G = \text{the sediment load (kg m^{-2} \cdot s^{-1}), and}$
- $T_c$  = the sediment transport capacity of surface runoff (kg m-2·s-1).





- Is the amount of sediment that the flow is capable of moving.
- The flow may not be capable of transporting all of the eroded material.
- The flow may be capable of transporting more than the eroded material.
  - In this case the transport capacity will not be satisfied.







- Transport capacity is for bulk particles in suspension. Particles will be transported in relation to fraction in suspension.
- Formulation:
- $q_s = sediment unit discharge (ton m<sup>-1</sup> s<sup>-1</sup>),$ 
  - q = unit water discharge (m<sup>2</sup> s<sup>-1</sup>), and
  - S<sub>f</sub> = friction slope.
  - The erodibility constant K is the product of the three factors

 $q_s = 25500q^{2.035} S_f^{1.664} \frac{K}{0.15}$ 

- Soil erodibility factor (0-1) large stones to fine silt.
- Cover factor (0-1) concrete to bare.
- Crop management factor (0-1) perfect to none.





#### **Engelund-Hansen**

Transport for individual particle sizes or types

$$G_{i} = K_{i} \frac{0.0 B5^{2} h^{3/2} S^{3/2}}{(s-1)^{2} D_{i} \sqrt{g}}$$

- *G<sub>i</sub>* = the volumetric sediment transport rate of *i*-th size fraction,
- K = the calibration coefficient (= 1 for standard equation),
- $F_i$  = the proportion of *i*-th faction in the parent material or deposited layer,
- B = the width of flow,
- V = the mean water velocity,
- h = the flow depth,
- S = the water surface slope,
- s = is the specific gravity of *i*-th fraction,
- g = the gravitational acceleration,
- *D<sub>i</sub>* = the mean size of *i*-th fraction.
- Appropriate method if particles are from materials other than quarts, i.e. the specific gravity is not approximately 2.65.





- Everaert (1991) conducted flume experiments to attempt to relate the D<sub>50</sub> (median grain size μm) and several physical constants to observed unit discharge of sediment (q<sub>s</sub> kg m<sup>-</sup> <sup>2</sup> s<sup>-1</sup>) in the flume.
- While he derived many such formulations, we found only four to produce usable results, as transport capacity is implemented in GSSHA.
- One advantage of using these methods is they require no parameter specifications. All parameters are calculated internally as follows.



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#### **Effective Stream Power**

- Stream Power (N m<sup>-1</sup> s<sup>-1</sup>)
- Density
- Gravity
- Unit Discharge
- Friction slope
- Effective Stream Power

$$w = \rho g q S_f$$

 $= 4.610^{-7} \Omega^{1.75} D_{50}$ 

$$\Omega = \frac{\omega^{1.5}}{R^{\frac{2}{3}}}$$

 $q_s$ 

Unit Discharge (kg m<sup>-1</sup> s<sup>-1</sup>)



- Unit Sediment Discharge
- (kg m<sup>-2</sup> s<sup>-1</sup>)

 $q_s = 0.316 \left(S_f Vel\right)^{2.59} D_{50}^{-0.39}$ 

D50 in  $\mu m$ 





#### **Slope and Unit Discharge**

■ <i>D50</i> <33	$q_s = 6.16595 (S_f)^{1.9} q^{1.1}$
■ 33 < <i>D50</i> < 61	$q_s = 10.964 \left(S_f\right)^{1.94} q^{1.1}$
■ 61 < <i>D50</i> < 122	$q_s = 9.332 \left(S_f\right)^{1.77} q^{1.79}$
■ 122< <i>D50</i> < 190	$q_s = 64.565 (S_f)^{2.96} q^{2.1}$
■ <i>D50</i> > 190	$q_s = 10.964 \left(S_f\right)^{1.94} q^{1.1}$

*D*50 in μm 00-0

2.18

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- Shear Stress (N m<sup>-2</sup>)
- Shear Velocity (m s<sup>-1</sup>)
- $D_{50<33 \,\mu>1.4} \qquad q_s = 0.035 \,(u 1.40)^{2.88}$
- 33<D50<61 μ>1.4

 $q_s = 0.052 \, (u - 1.40)^{2.95}$ 

 $\tau = \rho g h S_f$ 

 $u = \left(\frac{\tau}{\rho}\right)$ 

D50 in  $\mu m$ 

- 61<*D*50<122  $\mu$ >1.45  $q_s = 0.029 (u 1.45)^{3.74}$
- 122<*D*50<190  $\mu$ >1.55  $q_z = 0.024 (u-1.40)^{4.14}$

O.

μ > 1.8

 $q_s = 0.0092 (u - 1.80)^{5.06}$ 

20-



### **Shear Velocity**





**Computation of D50** 

 D50 is computed from the material properties in each overland cell. Each cell can/will have a unique distribution of particles and particle sizes that evolves over time as erosion and deposition occur. Assuming a semi-log distribution, the mean diameter is computed as (http://cirp.usace.army.mil/wiki/CMS-Flow\_Multiplesized\_Sediment\_Transport) :

$$d_k = \exp\left[\ln d_1 + \ln (d_N/d_1) \frac{k-1}{N-1}\right]$$



**Particle Settling** 

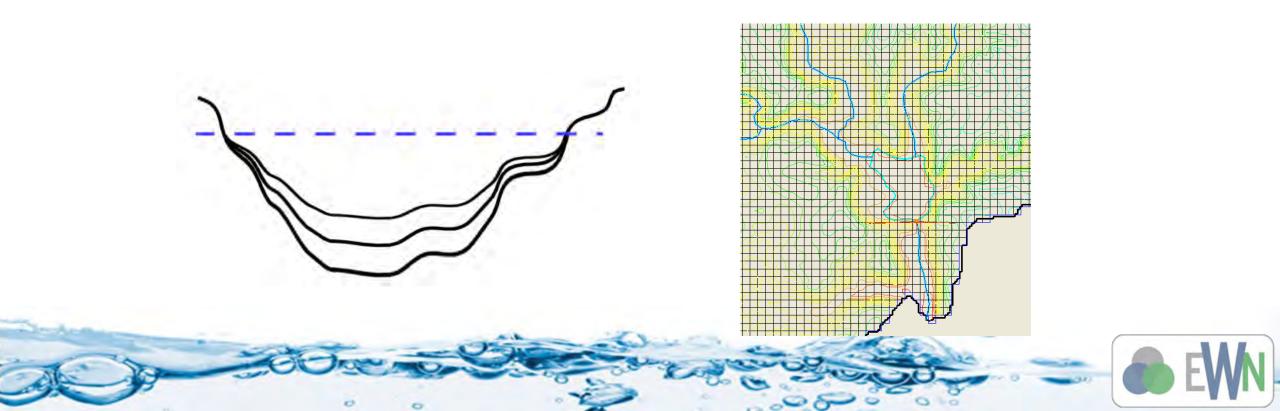
Particle settling is calculated using the trap efficiency (Johnson, 2000) which describes the fraction of each particle size detained in the current cell

$$T_{Ei} = 1 - e^{\frac{-X\omega_i}{hV}}$$

- $T_{Ei}$  is the trap efficiency for the i<sup>th</sup> size fraction,
- X is the grid cell size (m),
- $\omega_i$  is the fall velocity of the i<sup>th</sup> size fraction (m s<sup>-1</sup>),
- h is the overland flow depth (m), and
- V is the overland flow velocity (m s<sup>-1</sup>).



- Particles larger than sand (user specified) treated as bed load and routed with Yang's method
- Smaller particles treated as wash load advection dispersion equation
- Stream cross sections adjusted for erosion and deposition
- Particles settle or can be passed through reservoirs





- Sands from overland plane into reservoir are deposited in boundary reservoir cells.
- Sands from upstream channel network are removed from system and accounted for.
- Fines from overland and channels are accounted in reservoir, and can flow through.
- Reservoir is a completely mixed reactor.
- Fines settle in overland cells within the reservoir.
- Fines are discharged at the reservoir outlet.

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#### **Interaction with Reservoirs**

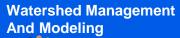
 $\begin{array}{l} \underline{\textbf{Terminal velocity}}\\ W = Weight \ of \ object\\ F_b = Buoyancy \quad force \end{array}$  $concentration[k] = \frac{suspended volume[k]}{reservoirs volume}$  $D^{\nu} = Drag$  force  $W = F_b + D$  $V_t[K] = \frac{gd^2}{18\mu(\rho_s - \rho)}$ 

settled  $vol[k] = concentration[k] V_t[k] dt$  reservoir area

The settled volume is uniformly distributed over all cells currently within the lake.

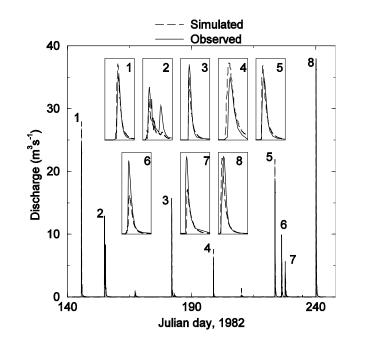
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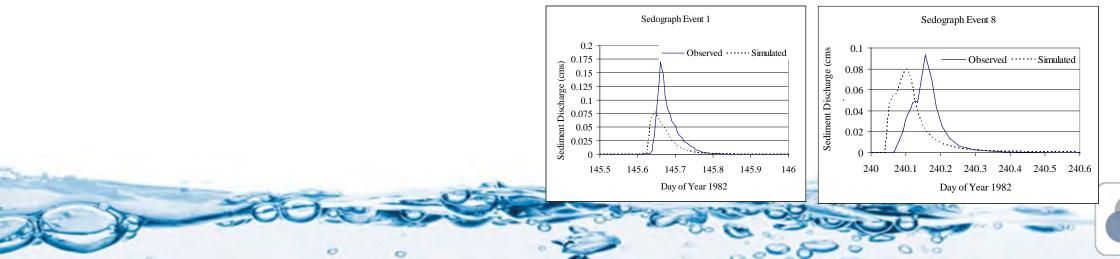




### **Continuous Simulations**

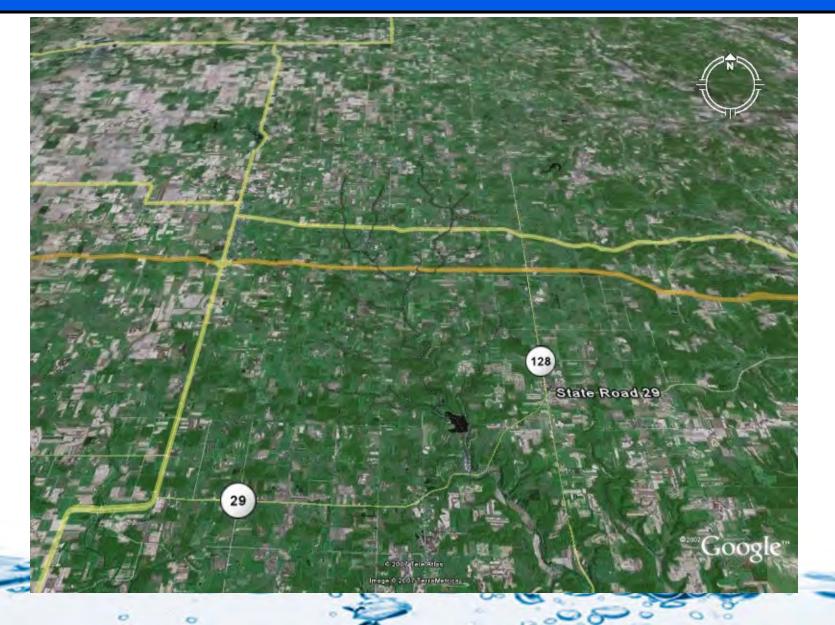
- Simulates each event with memory of previous events
  - Erosion
  - Deposition
  - Changes in particle distribution
- Simulates events more than 4 orders of magnitude different without change in parameter values





# And Modeling

#### Eau Galle Sediment Transport







#### **Simulating Sediment Transport**

- Select Sediment erosion in the job properties.
- Specify the transport capacity method.
- Specify the distributions of particle sizes and specific gravities.
- Specify the sediment erosion parameters.
  - Index map
  - Mapping table
- Specify the fraction of each particle size in each soil type.
  - Index map
  - Mapping table
- Set the stream erosion parameters.
  - Water temperature
  - Maximum erosion in each stream link
  - Particle size of sand
    - All particles sand size or larger will be treated as bed load
    - All particles smaller than sand size are treated as wash load.

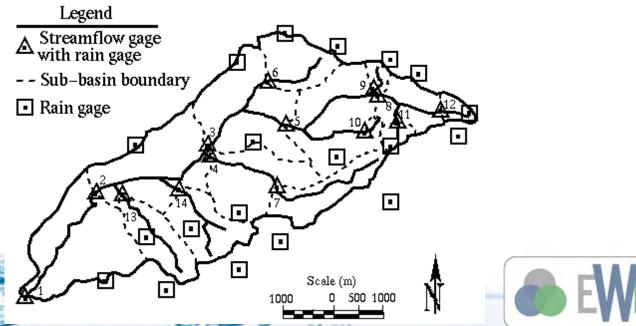


#### Goodwin Creek Experimental Watershed near Oxford, Mississippi, est. 1981.

#### Operated by:

US Dept. of Agriculture, Agricultural Research Service, National Sedimentation Laboratory



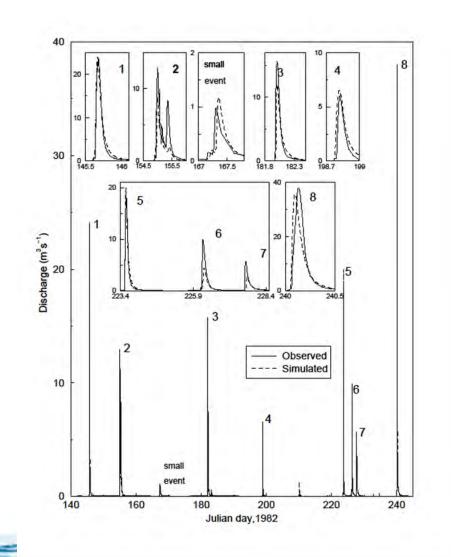


**Case Study** 



## **GSSHA Model Hydrologic Performance**

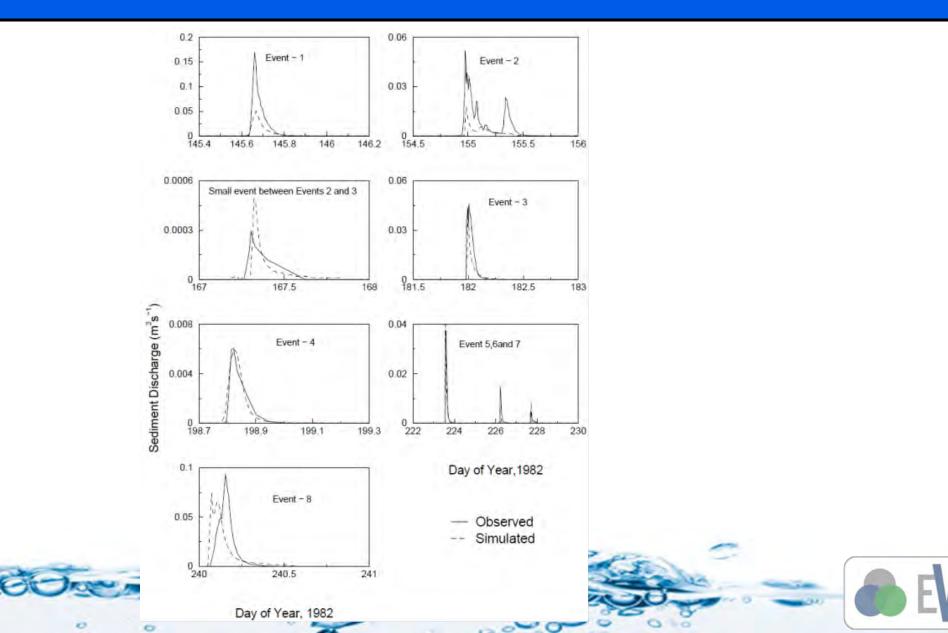
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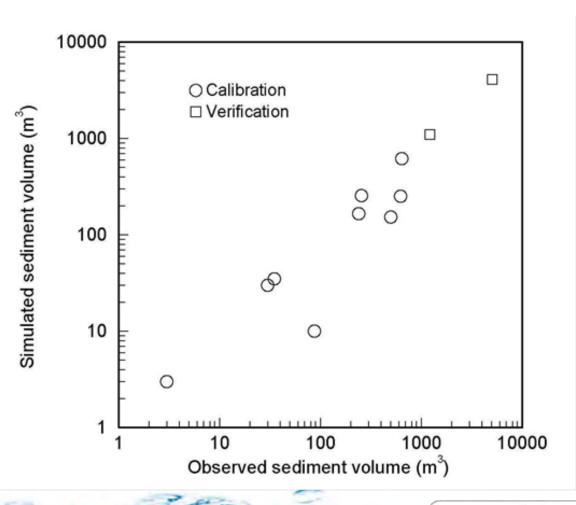
#### **Suspended Sediment Calibration/Verification Results**



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- Included two additional large events in verification period.
- Four orders of magnitude in event size.
- Mean absolute error across all events is 28%.
- No real trend between event size and accuracy.





- If the particle is not silica based, the specific gravity is not 2.65, you must use the England-Hansen equation.
- If you have observed sediment data to calibrate to, use the Kilinc-Richardson equation.
- If you have no observed sediment data to calibrate to, use the slope-unit discharge equation.







- Chapter 10 of the User's Manual on the GSSHA wiki.
- Downer, C. W., F. L. Ogden, N. R. Pradhan, S. Liu, and A. R. Byrd. 2010. Improved soil erosion and sediment transport in GSSHA. *ERDCTN-SWWRP-10-3*, U.S. Army Engineer Research and Development Center, Vicksburg, MS. <u>https://swwrp.usace.army.mil/</u>
- Downer, C. W., N. R. Pradhan, F. L. Ogden, and A. R. Byrd, 2014. Testing the effects of detachment limits and transport capacity formulation on sediment runoff predictions using the US Army Corps of Engineers GSSHA model. JHE, 04014082 1-11, doi: 10.1061/(ASCE)HE.1943-5584.0001104.